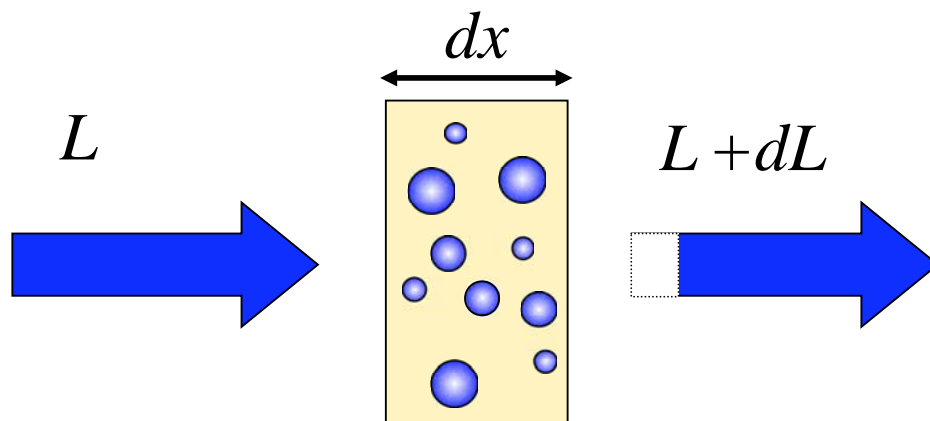


2. Radiation interaction with materials

Lambert-Beer's law: Differential form

- Monochromatic radiation (単色の光)
- (Volume) Extinction coefficient (消散係数) $[e] = 1/m$

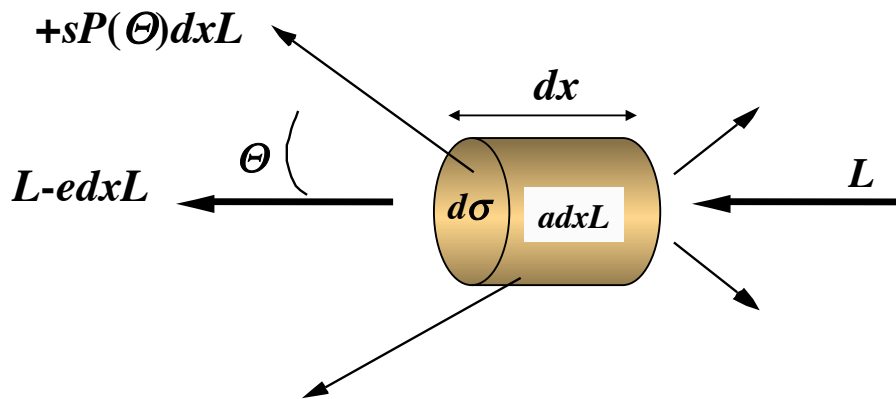
$$dL = -eLdx$$



Scattering phenomenon (散乱現象)

- Scattering of radiation to other angles
- Scattering angle (散乱角): Θ
- Scattering coefficient (散乱係数): s (m^{-1})
- Scattering phase function (散乱位相関数): $P(\Theta)$
- Direct radiance: $dL = -edxLd\Omega$
- Scattered radiance: $dL = sP(\Theta)dxLd\Omega$

$$\int_{4\pi} P(\Theta)d\Omega = 2\pi \int_{-1}^1 P(\Theta)d \cos \Theta = 1$$



Absorption and scattering

- Absorption and scattering phenomena with coefficients for absorption and scattering: a ($1/\text{m}$) and s ($1/\text{m}$)
- Energy conservation: $e Ldx = a Ldx + s L dx$ leads $e = a + s$
- Single scattering albedo: $\omega = s/e$
- Particle system composed of (equivalent) spheres with radius r with a number density N ($1/\text{m}^3$)
- Geometrical cross section: C_{geo}
- Optical cross sections (coefficients) for absorption, scattering and extinction: $C_{abs}, C_{sca}, C_{ext}$
- Efficiency factors: only depend on the size parameter $\alpha = kr$

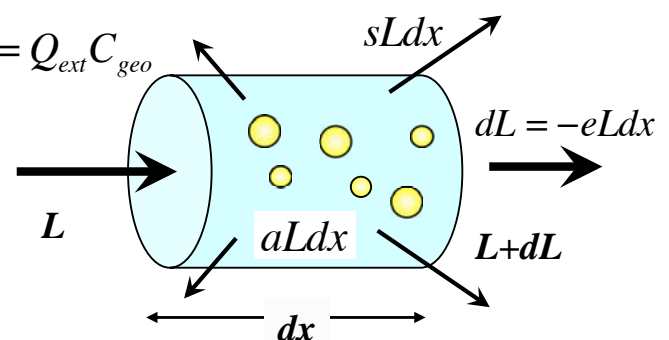
$$C_{geo} = \pi r^2$$

$$C_{abs} = Q_{abs} C_{geo}, \quad C_{sca} = Q_{sca} C_{geo}, \quad C_{ext} = Q_{ext} C_{geo}$$

$$C_{ext} = C_{abs} + C_{sca}$$

$$\alpha = kr = \frac{2\pi r}{\lambda}$$

$$a = C_{abs} N, \quad s = C_{sca} N, \quad e = C_{ext} N$$



Lambert-Beer's law: Integral form

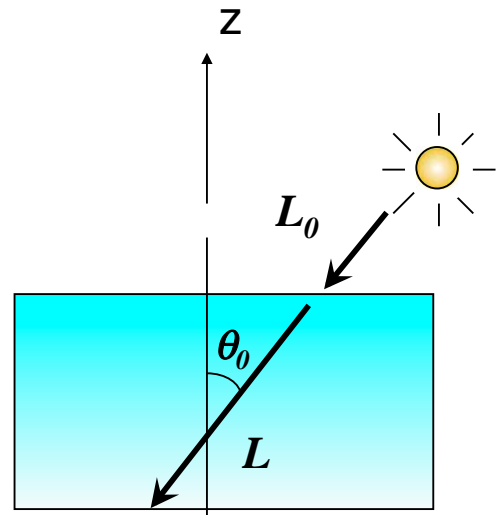
- Direct radiation (直達放射)
- Optical thickness (光学的厚さ), depth $[m] \times [1/m] = [1]$
- Optical airmass m
- Solar zenith angle (太陽天頂角)

$$dL = -eLdz / \cos \theta_0$$

$$L = L_0 \exp(-\tau m)$$

$$T \equiv L / L_0 = \exp(-\tau m)$$

$$\tau = \int_z^{\infty} e(z) dz, \quad m = 1 / \cos \theta_0$$



Large particle limit

- Large particle limit
 - Refraction and diffraction

$$Q_{ext} \rightarrow 2,$$

$$Q_{sca} \rightarrow 2 (\omega = 1), \quad \rightarrow 1 (\omega < 1)$$

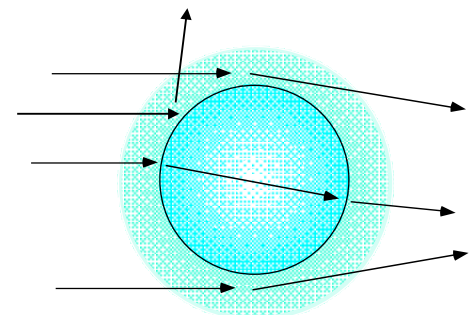
$$e = Q_{ext} C_{geo} N = 2\pi r^2 N$$

$$\tau = e \Delta x$$

$$T = \exp(-m\tau)$$

- But, why is the cloudy sky so bright?

Ray optics (幾何光学)



r (micron)	10
N (particles/cc)	100
e (/m)	0.0628
dx (m)	100
tau	6.28
Solar zenith angle (deg)	60
Optical airmass	2.00
Direct transmittance	3.55E-06

Refractive index

- Complex refractive index (複素屈折率)
- Snell's law
- Lambert absorption coefficient

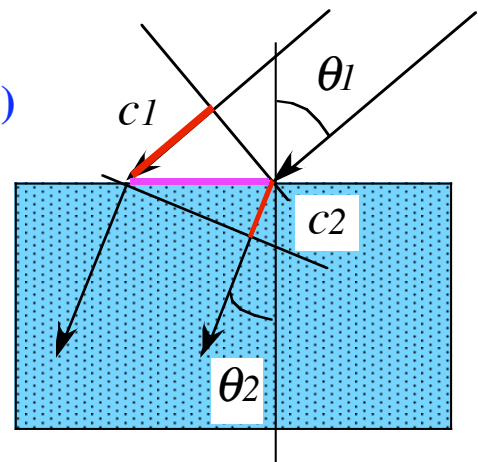
$$E = E_0 e^{i(\omega t - \tilde{m}k_0 x)}$$

$$\tilde{m} = m_r - im_i$$

$$\frac{\sin \theta_1}{\sin \theta_2} = \frac{c_1}{c_2} = \tilde{m}_{12}, \quad \omega = k_1 c_1 = k_2 c_2, \quad k_2 = \tilde{m}_{12} k_1$$

$$|E|^2 = |E_0 e^{i(\omega t - \tilde{m}k_0 x)}|^2 = |E_0|^2 e^{-\Lambda x}$$

$$\Lambda = 2m_i k_0 = \frac{4\pi m_i}{\lambda}$$



wavelength	0.5 micron
mi	-0.01
Lambert abs	-251.2 for 1 mm thickness

Small particle limit

Rayleigh (1871)
van de Hulst (1957)

- Small dielectrics; Rayleigh scattering (Dipole scattering): $r^6 \lambda^{-4}$
- Loosing the efficiency of scattering
- Infrared region: Absorption is dominant

$$\alpha = ka = \frac{2\pi a}{\lambda}$$

$$Q_{ext} = \varepsilon_1 \alpha + \varepsilon_3 \alpha^3 + \varepsilon_4 \alpha^4$$

$$Q_{sca} = s_4 \alpha^4, \quad \text{if } \alpha < 1.5$$

$$s \approx s_4 \alpha^4 \pi r^2 N = s_4 \pi (2\pi)^4 r^6 \lambda^{-4} N$$

$$m_i = 0: \quad \varepsilon_4 = s_4, \quad \omega = 1$$

$$m_i > 0: \quad \omega \rightarrow 0, \quad e \propto V$$

$$\alpha = (m_r^2 + m_i^2)^2, \quad \beta = m_r^2 - m_i^2, \quad \gamma = m_r m_i$$

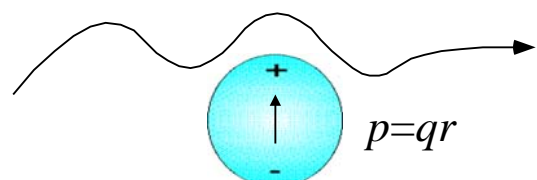
$$z_1 = \alpha + 4\beta + 4, \quad z_2 = 4\alpha + 12\beta + 9$$

$$\varepsilon_1 = \frac{24\gamma}{z_1}, \quad \varepsilon_3 = \gamma \left[\frac{4}{15} + \frac{20}{3z_2} + \frac{24}{5z_1^2} (7\alpha + 4\beta - 20) \right]$$

$$\varepsilon_4 = \frac{8}{3z_1^2} [(\alpha + \beta - 2)^2 - 36\gamma^2]$$

$$s_4 = \frac{8}{3z_1^2} [(\alpha + \beta - 2)^2 + 36\gamma^2]$$

もうひとつの導出方法
Dipole scattering
(双極子散乱)



Small particle limit (2)

$$m_r = 1 + \Delta m, \quad m_i = 0$$

$$\alpha = m_r^4 = 1 + 4\Delta m, \quad \beta = m_r^2 = 1 + 2\Delta m, \quad \gamma = 0$$

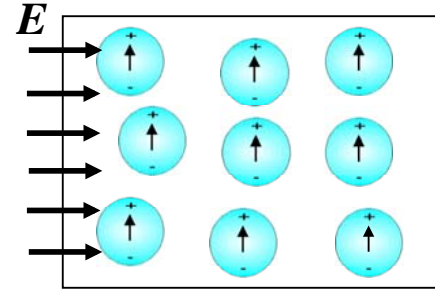
$$z_1 = 1 + 4\delta + 4 + 8\delta + 4 = 9 + 12\delta \approx 9$$

$$s_4 = \frac{8}{3z_1^2} (\alpha + \beta - 2)^2 = \frac{2^5}{3^3} \Delta m^2$$

$$C_{sca} \approx s_4 x^4 \pi r^2 = \frac{2^5}{3^3} \Delta m^2 \pi (2\pi)^4 \lambda^{-4} r^6$$

$$N = 1 / \left(\frac{4\pi}{3} r^3 \right)$$

$$C_{sca} = \frac{2^5}{3^3} \Delta m^2 \pi (2\pi)^4 \lambda^{-4} \left(\frac{3}{4\pi N} \right)^2 = \frac{32\pi^3}{3\lambda^4} \left(\frac{\Delta m}{N} \right)^2$$



Rayleigh scattering

- **Refractive index of air** $m - 1 = 2.9E-4(1atm, 0C)$

$$C_{sca} \approx \frac{32\pi^3}{3\lambda^4} \left(\frac{m-1}{N} \right)^2 = \frac{3.84E-32}{\lambda_{micron}^4} (m^2)$$

$$N = N_a n / V = N_a P / RT = 6.02E23 * 1.013E5 / 8.31 / 273 = 2.69E25 (/m^3)$$

- **Column air molecule number**

$$N_{air} = 1.013E5 (N / m^2) / 9.8 (m / sec^2) * 6.02E23 / 0.029 = 2.1E29 (/m^2)$$

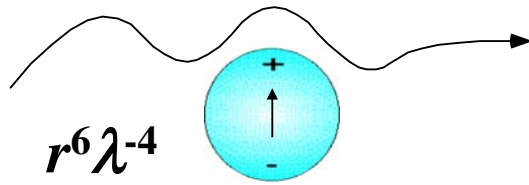
- **Rayleigh optical thickness of the atmosphere**

$$\tau_R = N_{air} C_{sca} = \frac{32\pi^3 N_{air}}{3\lambda^4} \left(\frac{m-1}{N_{1atm}} \right)^2 = \frac{0.0081}{\lambda(\mu m)^4}$$

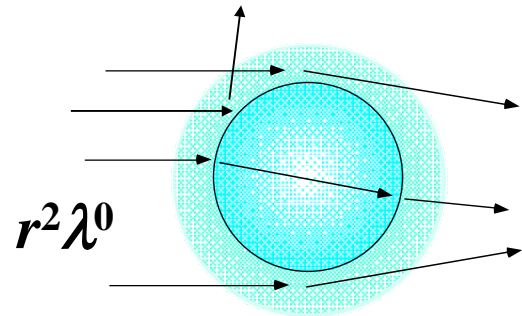
$$\tau_R = 0.00864 \lambda^{-(3.916+0.074\lambda+0.050/\lambda)} P_{atm}$$

Multipole moments by dielectrics

Small particle limit:
Rayleigh scattering

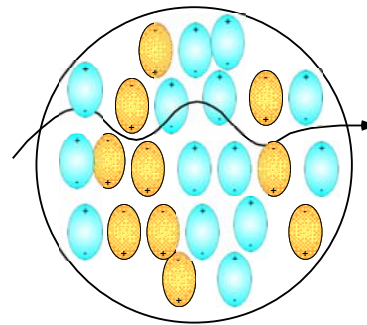


Large particle limit:
Geometrical optics



Large dielectric sphere:
Mie scattering theory (1908)

100 year anniversary in 2008!

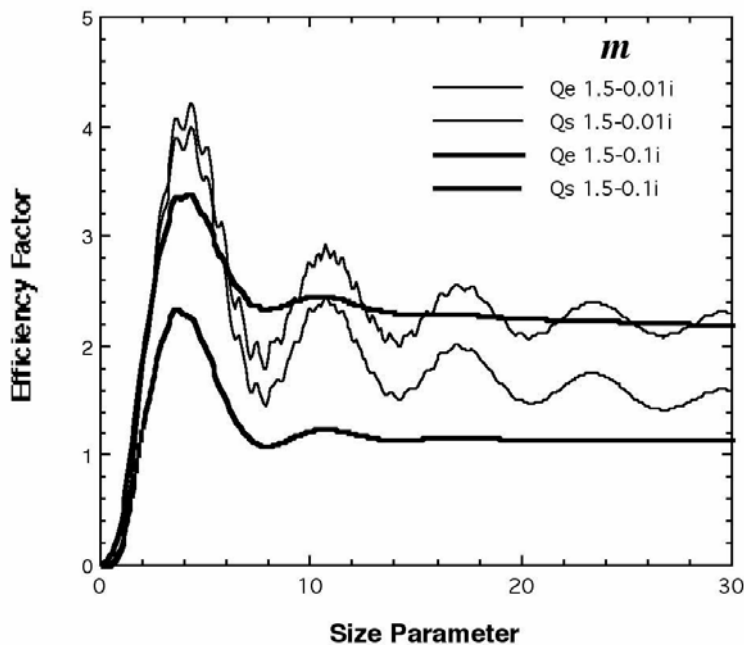


Efficiency factors

- Efficiency factors
- Critical size parameter

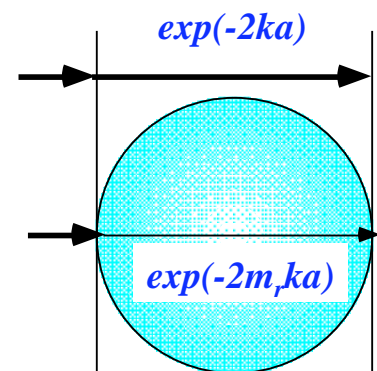
$$\alpha = kr = \frac{2\pi r}{\lambda}$$

$$C_{ext} = Q_{ext} \pi r^2, \quad C_{abs} = Q_{abs} \pi r^2, \quad C_{sca} = Q_{sca} \pi r^2$$

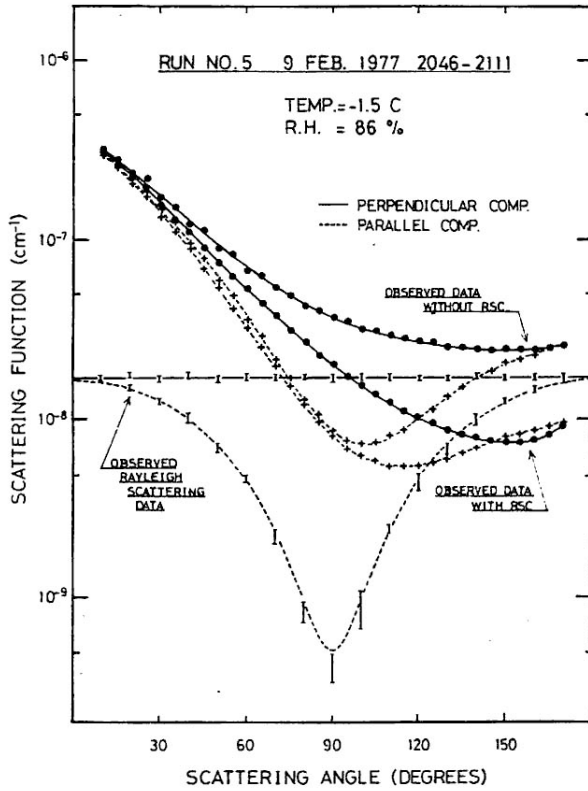


$$\alpha_c = \frac{\pi}{2(m_r - 1)}$$

$$2m_r ak - 2ak = \pi$$



Contribution of atmospheric molecules



- Rayleigh scattering
- Gaseous absorption (water vapor, Ozone)

$$\tau_R = 0.00864 \lambda^{-(3.916+0.074 \lambda+0.050/\lambda)} P_{atm}$$

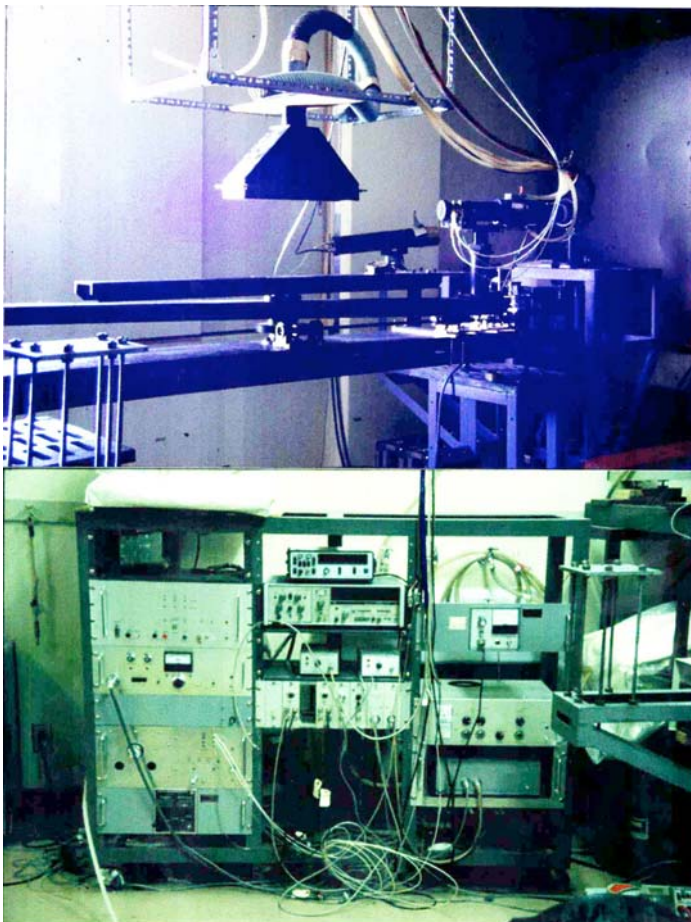
$$\tau = \tau_{aerosol} + \tau_{molecule}$$

$$\omega \tau = \omega_a \tau_a + \omega_m \tau_m$$

$$\omega \tau P(\Theta) = \omega_a \tau_a P_a(\Theta) + \omega_m \tau_m P_m(\Theta)$$

$$P_m(\Theta) \approx \frac{3}{16\pi} (1 + \cos^2 \Theta)$$

Tanaka et al. (1983)



Polar nephelometer

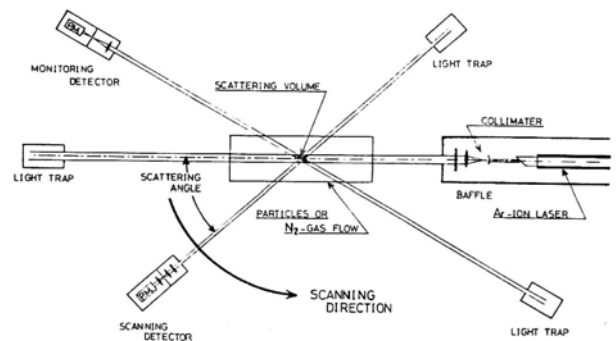
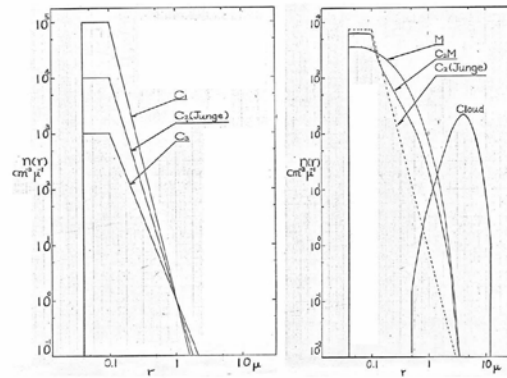


Fig. 1 Schematic diagram of polar-nephelometer.

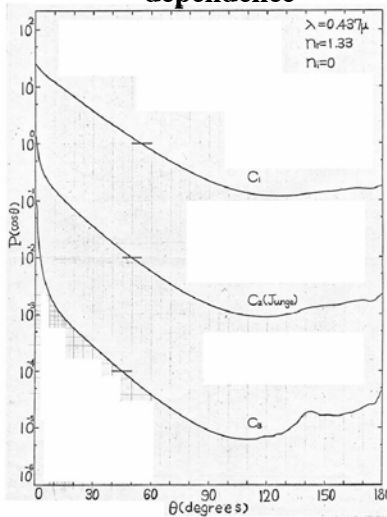
Takamura and Tanaka (1978)

Phase function characteristics

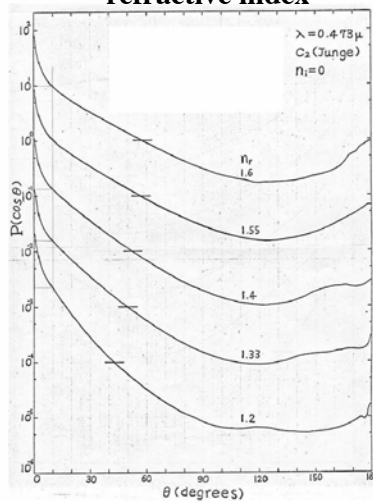
K. Sato (1973)



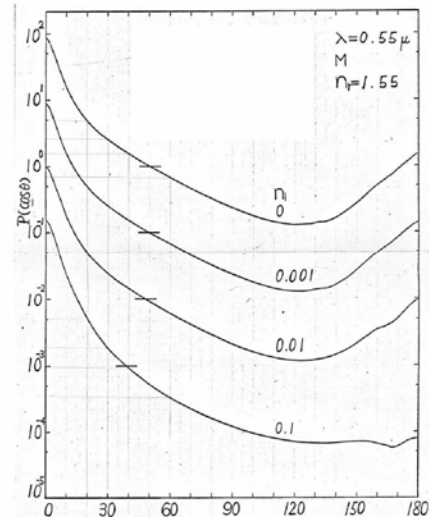
Size distribution dependence



Real part of refractive index



Imaginary part



Particle size distribution (粒径分布)

- Size distribution, Size spectrum

$$v(\ln r) d \ln r = \frac{4\pi r^3}{3} n(r) dr, \quad v(\ln r) = \frac{4\pi r^4}{3} n(r)$$

- Power law size distribution

$$n(r) = \begin{cases} Cr^{-p} & \text{if } r > r_0 \\ Cr_0^{-p} & \text{if } r \leq r_0 \end{cases}$$

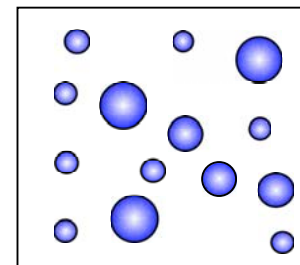
➤ Junge distribution: $p=4$

- Log-normal distribution

$$n(r) = \sum_n \frac{N_n}{\sqrt{2\pi}\sigma_n} \exp\left\{-\frac{1}{2}\left(\frac{\ln r - \ln r_{m,n}}{\sigma_n}\right)^2\right\}$$

- Modified Gamma distribution

$$n(r) = Cr^\alpha e^{-\beta r^\gamma}$$



Observed size distributions

Pasceri, R.E., and S.K. Friedlander (1965)

$$N(r) = \int_r^{\infty} n(r) dr$$

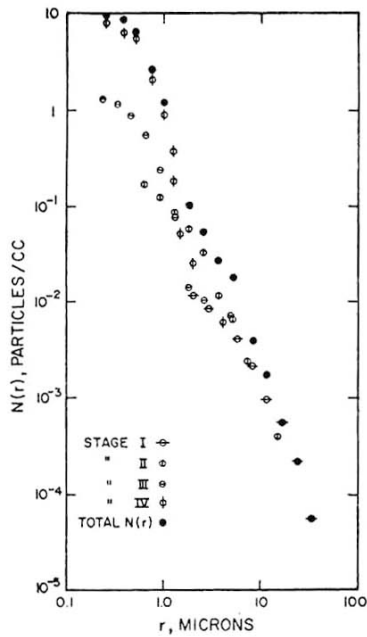


FIG. 1. Run 1 made outdoors in Baltimore 14 February 1961. Cumulative data for individual impactor stages are combined to give cumulative distribution for the run.

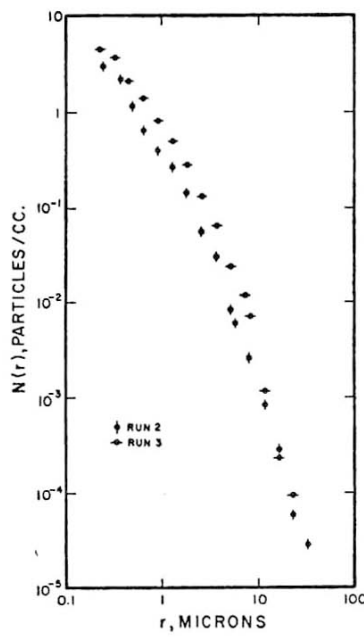


FIG. 2. Runs 2 and 3 made outdoors in a residential area outside Baltimore 30 June 1961 and 1 July 1961. The results are close to those shown in Fig. 1.

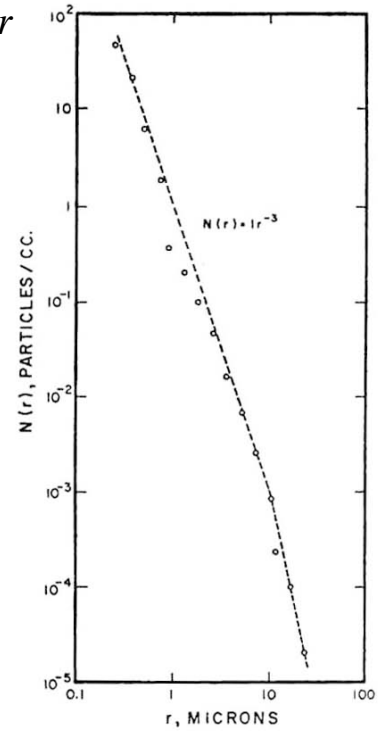
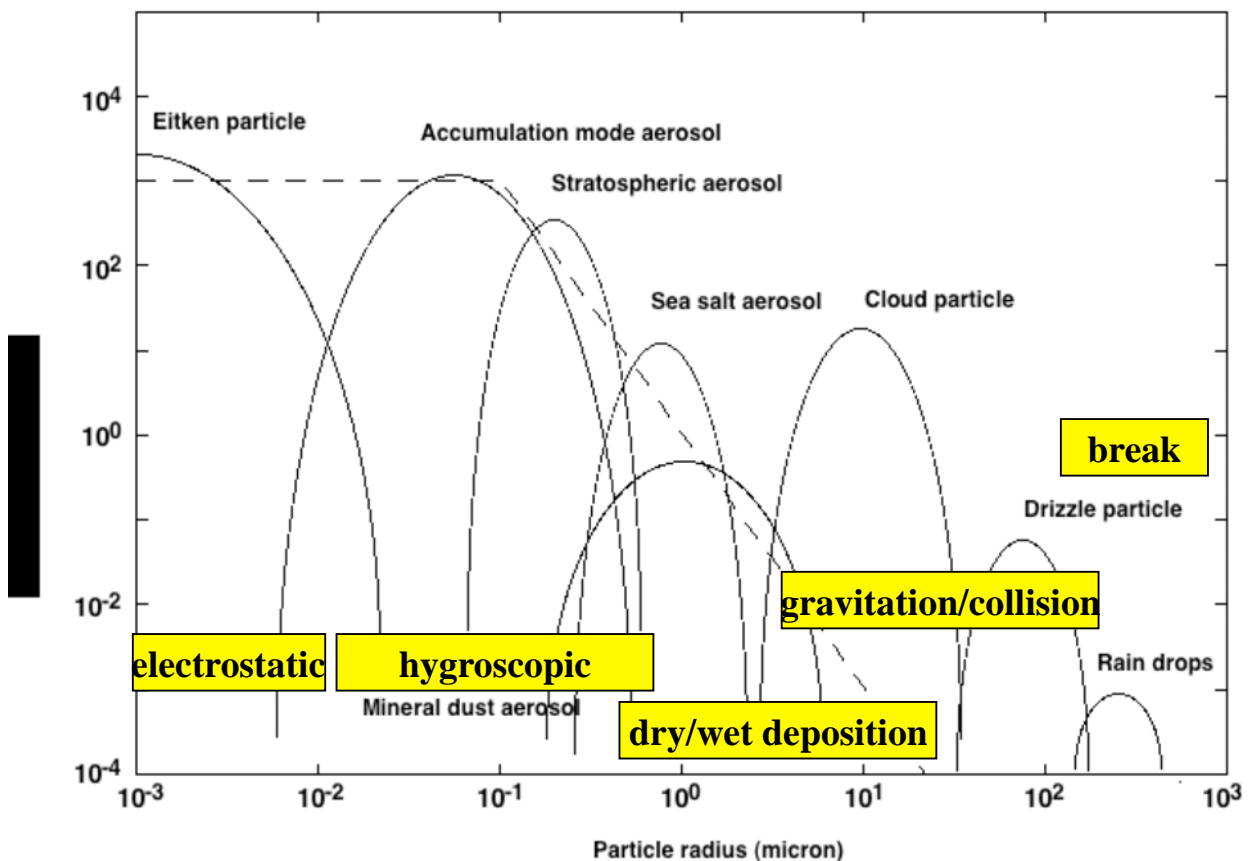


FIG. 3. Run 4 made in the laboratory on 22 February 1962. Over the radius range 0.3 to 10 microns the data are well represented by $N(r) \sim r^{-3}$.

Size distribution in the atmosphere



Ångström's law

- Power law size distribution $n(r) = \begin{cases} Cr^{-p} & \text{if } r > r_0 \\ Cr_0^{-p} & \text{if } r \leq r_0 \end{cases}$
- p=4 : Junge
- Aerosol distribution tends to have a power law type size distribution, then

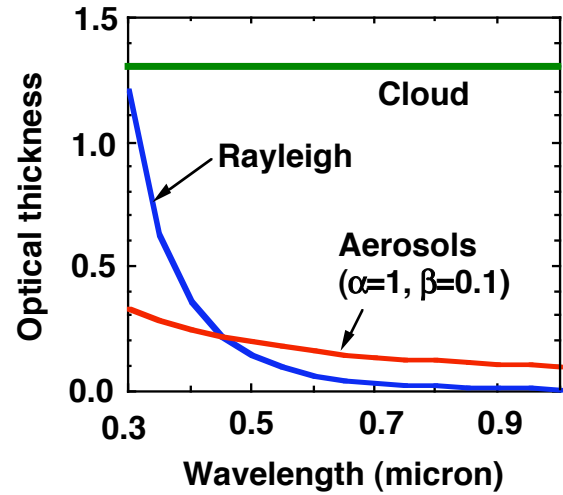
$$e = \int_0^\infty \pi r^2 Q(x) Cr^{-p} dr = Ck^{p-3} \int_0^\infty \pi x^2 Q(x) x^{-p} dx$$

- Ångström's law

$$\tau = \beta \lambda^{-\alpha} \quad \alpha = p - 3$$

- Junge distribution

$$\alpha = 1$$



Size distribution in the ocean (1)

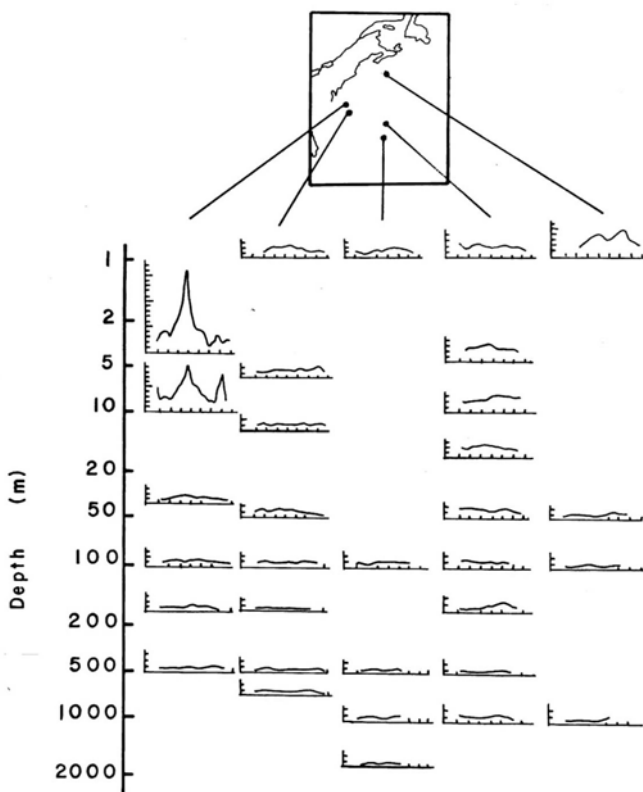


FIG. 8. Size distributions of suspended particulate matter at various depths in the western North Atlantic.

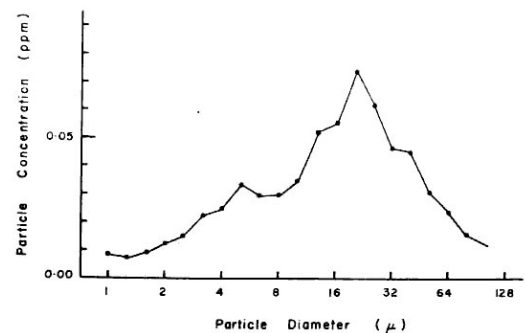
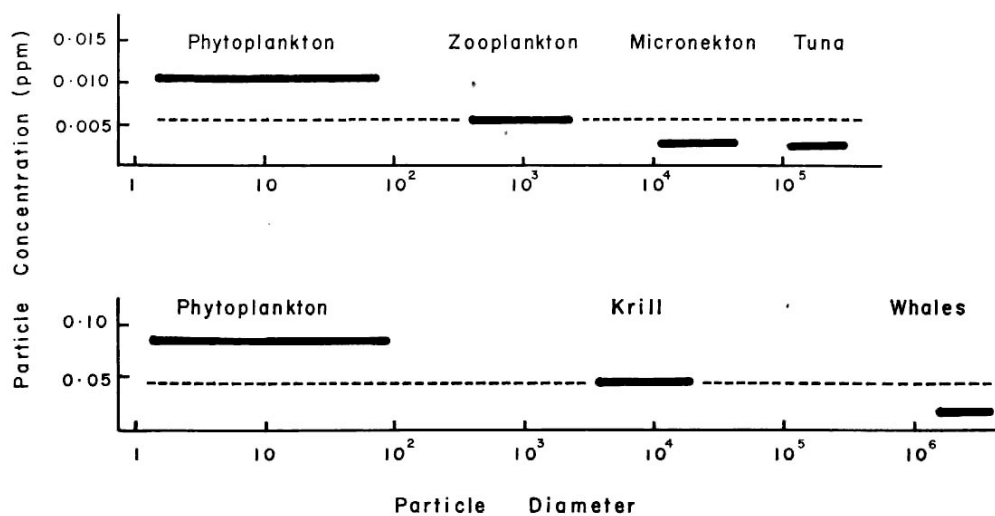


FIG. 2. Size-frequency distribution; to show the notation of the axes and the least number of data points used to define the form of the distribution. Concentration is *by volume*. All the size-frequency distributions in Figs. 3-9 were constructed in this way.

Size distribution in the ocean (2)



$$\frac{dV}{d \ln r} = \frac{4\pi r^3}{3} \frac{dN}{dr} \frac{dr}{d \ln r} = \frac{4\pi r^4}{3} n(r)$$

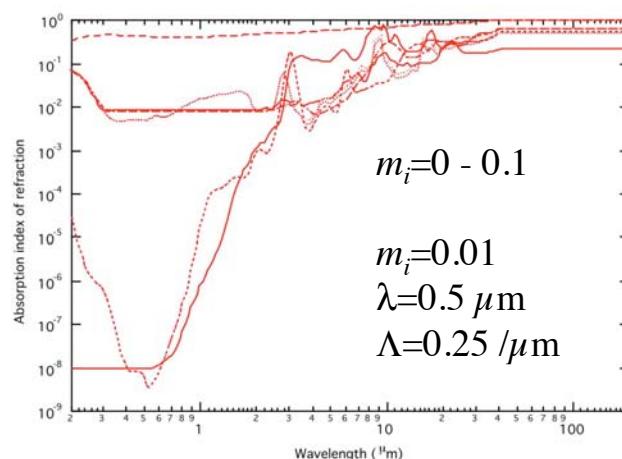
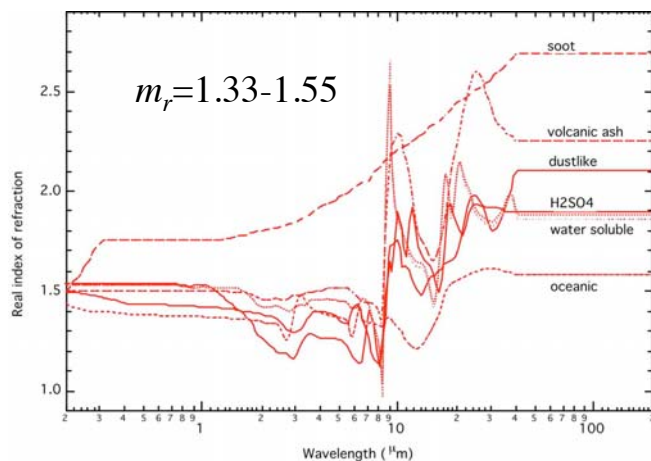
Sheldon et al. (1972)

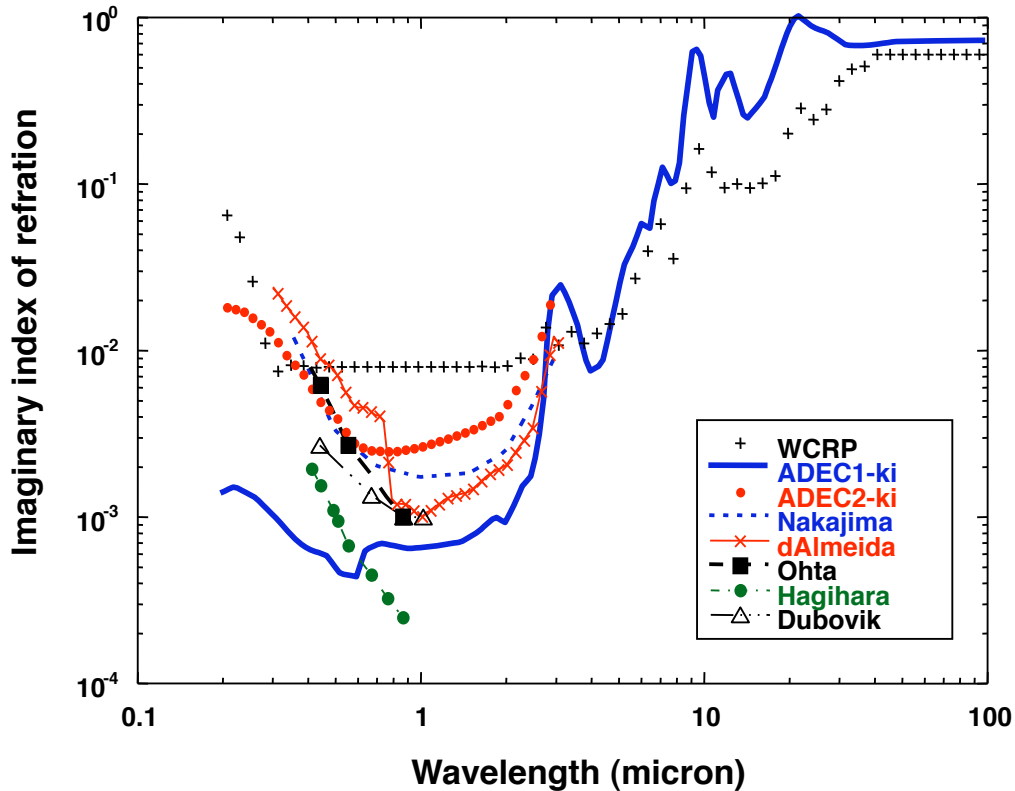
Refractive index of aerosols

- Particle types

- Water, ice
- Soot, volcanic ash, dust like, H₂SO₄, water soluble, oceanic

$$|E|^2 = |E_0 e^{i(\omega t - \tilde{m} k_0 x)}|^2 = |E_0|^2 e^{-\Lambda x}, \quad \Lambda = 2m_i k_0 = \frac{4\pi m_i}{\lambda}$$





Nakajima et al. (2007)

Einstein-Smolukowski theory (1908, 1910)

$M_a=29$ g/mole, $M_w=18$ g/mole

$m(\text{air atm}/\text{m}^3) = 1000 \text{ litre} / 22.4 * 29 \text{E}-3 = 1.29 \text{ kg}$

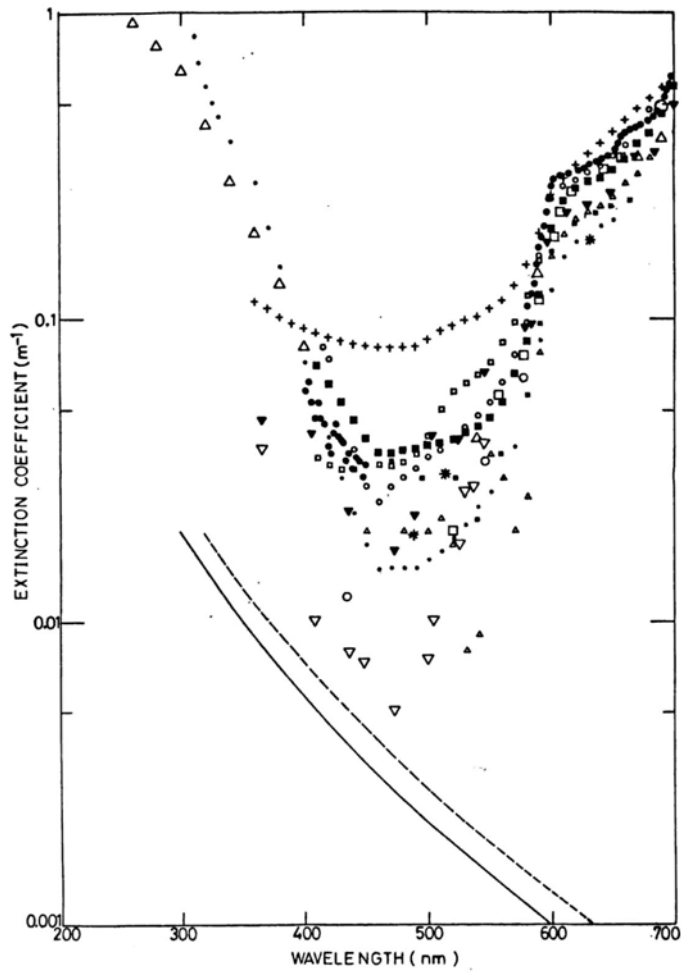
$m(\text{water}/\text{m}^3) = 1000 \text{ kg}$

$N_{\text{water}}/N_{\text{air}} = (m(\text{water})/M_w) / (m(\text{air})/M_a) = 1250$

$$C_{s,\text{water}} N_{\text{water}} \approx 1250 C_{s,\text{air}} N_{\text{air}} = 0.0013 \lambda_{\text{micron}}^{-4} \quad (1/m) \quad ???$$

Observed value for water:

$$C_{sca} = (1.934 + 0.9017 \lambda^{-4.5}) \times 10^{-4} \quad (1/m)$$



Extinction and scattering coefficients of pure water